Geochemistry of rare earth elements and yttrium in a Ge-poor coal from the Wulantuga ore deposit, Inner Mongolia, North China

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Abstract This paper reports data of yttrium and rare earth (REY, or REE if Y is considered) in a Ge-poor (the average value of Ge is $0.57 \mu g/g$ coal from the Wulantuga ore deposit, middle Inner Mongolia, northern China. The contents of yttrium and rare earth in six coal samples of the Wulantuga ore deposit were detected by inductively coupled plasma mass spectrometry. The content of yttrium and rare earth in the study area varies from 5.88 to 119.19 μ g/g, with an average of 61.85 lg/g, similar to the average value for world coals. The light-REY and heavy-REY enrichments are the two major REY-enrichment types of the Ge-poor coal from the Wulantuga ore deposit, and the top and bottom parts of the coal seam have a higher content of REY than the middle part. The sources of yttrium and rare earth in the Ge-poor coal of Wulantuga may be mainly associated with terrigenous materials and natural waters enriched in yttrium and rare earth.

Keywords Rare earth and yttrium · Wulantuga · Distribution pattern · Ge-poor coal

1 Introduction

Rare earth elements (REE) include La, Ce, Pr, Nd, Sm, Eu, Gd, Tb, Dy, Ho, Er, Tm, Yb and Lu, with atomic number from 57 to 71. Because of both similar geochemical characteristics and ionic radius to lanthanides, yttrium is generally considered as a rare earth element (Liu and Cao [1984\)](#page-4-0).

Yttrium and rare earth can be used as geochemical indicators in coal and coal-bearing strata because they own relatively stable geochemical characteristics during different geological processes (Li et al. [2005\)](#page-4-0). Yttrium and rare earth may provide significant information for the study of terrigenous areas and oxidation–reduction environment of

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peat accumulation, history of tectonic evolution, volcanic ash eruption, and hydrothermal solution activities. For instance, based on the analysis of modes of trace-element (including yttrium and rare earth) and mineral occurrence, Dai et al. [\(2014a](#page-4-0)) revealed the dominant sediment-source regions for the late Permian coal-bearing strata of the Huayingshan coalfield, eastern Sichuan, southwestern China, are the Dabashan and Leshan-Longnvsi Uplifts and the Hannan Upland. Moreover, the first finding of coal beds with high yttrium and rare earth elements contents in one of the coal-hosted ore deposits in Russia Far East has made the extraction of rare earth and Yttrium possible. Seredin [\(1996](#page-4-0)) noted that a total of several coal seams with similar and even higher contents (up to 1.0 % in ash) of yttrium and rare earth were found in six coal-bearing basins of this region. Owing to the increasing significance of REY role in industry, more and more attention is focused on the yttrium and rare earth in coal and coal combustion products.

Zhang et al. ([2008\)](#page-4-0) and Dai et al. [\(2014b](#page-4-0)) reported the concentrations of yttrium and rare earth in the Ge-elevated coal $(273 \mu g/g)$ from the Wulantuga deposit. However, data of yttrium and rare earth in the Ge-poor coal of Wulantuga are absent. This paper reports the data of six Gepoor coal samples from the coal-hosted Ge ore deposit from Wulantuga and thus could shed light on the yttrium and rare earth formation in the coal.

2 Geological background

The present study area is situated on the edge of southwestern Shengli coalfield, which was formed in the faultcontrolled Shengli sub-basin of the Erlian Basin in Inner Mongolia. The Shengli coalfield is in the Wunite depression in the west of Erlian Basin, and is located in a wide syncline (Huang et al. [2007\)](#page-4-0), with a 45-km length and 7.6 km width, and a surface area of 342 km^2 .

The Silurian, Devonian, Permian, Late Jurassic, Early Cretaceous, Neogene, and Quaternary are the sedimentary sequences in the Shengli coalfield (Fig. 1; Dai et al. [2012a](#page-4-0)). The Silurian and Devonian sedimentary strata have an average thickness of 2,363 m and are composed of various quartz schists. The minimum and maximum thickness of the Permian strata is 680 and 3,550 m, respectively. The Permian strata are covered by sandy mudstone, conglomerate, siltstone interbeded with limestone layers, lenses of andesitic tuffs and limestones. The lower part is made up of sandstone, andesite, interlayered with andesitic tuff, bioclastic rocks, limestone, marl, volcaniclastic rocks, mudstone and rhyolite. The upper Jurassic strata have a thickness of more than 3,734 m and consist of mafic and intermediate-felsic volcanic rocks (e.g., basalts, tuffs, volcanic breccia, quartz trachyandesite, and rhyolite). The coal-containing lower Cretaceous strata, collectively titled ''Bayanhua Group'', disconformably overlie the Jurassic system. It is made up of Hadatu, Aershan, and Saihantala Formations (Sha [2007](#page-4-0)). However the Hadatu Formation is absent in the area of the present study. As previously reported, the Aershan Formation of Gecontaining part of the Shengli Coalfield, is named the Tenggeer Group (Qi et al. [2007](#page-4-0)) or Xilin Formation, and the Saihantala Formation is named the Shengli Formation. The lower part with thickness from 70 to 564 m of the Xilin Formation is made up of fine sandstone, conglomerate, sandstone, and gravels interlayered with thin benches of mudstone. The middle part (generally 180–200 m) of the Xilin Formation is composed of coarse sandstone, grey mudstone, siltstone, and coal beds. The upper part (132.9–466 m, generally 160–180 m) of the Xilin Formation is mainly dominated by siltstone and mudstone. The No. 11 and 12 coal seams, with average thicknesses of 1.54 and 2.30 m respectively, are the two major coal beds in the Xilin Formation. The Neogene strata (average thickness 49.84 m) are predominantly composed of clay and gravels and disconformably overlie the Shengli Formation. It is noted that the distance between the early Cretaceous No. 6

Fig. 1 Coal-containing sedimentary strata of the Wulantuga ore deposit (after Dai et al. [2012a\)](#page-4-0)

coal collected in the deposit and the previously studied Gerich coal deposit is around 200 m.

3 Analytical procedures

A total of 6 coal ply samples were collected from the mining face of the mined No. 6 coal, which have an average thickness of 16.1 m at the Wulantuga Ge-poor deposit of the Shengli Coalfield in Inner Mongolia, North China. The collection method was following the Chinese Standard Method GB/T 482–2008 [\(2008](#page-4-0)). From top to bottom, the 6 coal ply samples were identified as WLTG-1 to WLTG-6.

Yttrium and rare earth in the coal samples were determined by ICP/MS (X series IIICP/MS). Before

geochemical analysis, the six coal samples were ground to less than 200 mesh and were digested using an UltraClave microwave high pressure reactor (Milestone). The mixedsolution of 330-mL distilled H₂O, 30-mL 30 % H₂O₂, and 2-mL 98 % H_2SO_4 were put into the digestion tank, served as the basic load. Initial nitrogen pressure for the wave digestion was 50 bars and the highest digestion-temperature was $240 \degree C$. Such high-pressure and high-temperature digestion process lasts for about 75 min. The reagents for 50-mg coal sample digestion consist of 2 mL 40 % HF and 5 mL 65 % HNO₃. Multi-element standards of Inorganic Ventures (CCS-1) were used for calibration of yttrium and rare earth content in tested samples as designed by Dai et al. [\(2011](#page-4-0)).

4 Results and discussion

As indicated in Table 1, the content of yttrium and rare earth in study area varies from 5.88 to 119.19 μ g/g, with an average of $61.85 \mu g/g$, similar to the average values for the world low-rank coals reported by Ketris and Yudovich [\(2009](#page-4-0)) (68.5 μ g/g). Table 1 illustrates that the lignite in the Wulantuga ore deposit is mainly enriched in light and heavy REY, and only sample WLTG-4 is enriched in medium REY. The top and bottom parts of the coal seam have higher contents of REY than that in the lower part. The δ Eu values vary from 0.53 to 1, with an average of 0.79, showing a pronounced Eu minimum. The δ Ce values vary from 0.75 to 0.99 (with an average of 0.92), indicating a slightly positive anomaly.

Content coefficients (REY content in the studied coal samples vs. world coals and Chinese coals) are presented in Fig. 2. With exceptions of La, Sm, and Er, the contents of other REY are lower than the average values for world and Chinese coals, respectively reported by Ketris and Yudovich ([2009\)](#page-4-0), and Dai et al. ([2012b\)](#page-4-0).

Fig. 2 Concentration coefficients of REY in Ge-poor coals of Wulantuga deposit. CC-C/C, concentration coefficients of coal samples versus Chinese coals; CC-C/W, concentration coefficients of coal samples versus world coals

Rare earth elements are generally classified into two types: Light-REE (LREE: La, Ce, Pr, Nd, Sm, and Eu) and Heavy-REE (HREE: Gd, Tb, Dy, Ho, Er, Tm, Yb, and Lu) (Han and Ma [2003\)](#page-4-0). Compared to the two-fold classification, the three-fold geochemistry classification of REY proposed by Seredin and Dai [\(2012](#page-4-0)), however, is more distinct to describe the REY distribution in coals and conventional REY ore deposits. This paper used the latter classification, and accordingly, normalized the original content of REY in coal samples by the abundance of REY in the upper continental crust (UCC; Taylor and McLennan [1985](#page-4-0)). The average REY contents of Chinese

Table 1 REY in the No. 6 coal from the Wulantuga ore deposit, Inner Mongolia (μ g/g)

Sample ID	REY	LREY	MREY	HREY	(La/Lu) _N	(La/Sm) _N	$(Gd/Lu)_{N}$	Enrichment type	δ Eu	δ Ce
WLTG-1	196.4	162.02	29.15	5.24	1.42	0.9	0.7		0.7	0.75
WLTG-2	17.33	9.83	6.4	1.11	0.39	0.58		H		0.9
WLTG-3	20.65	14.78	4.8	1.07	0.52	0.73	0.9	H	0.53	0.94
WLTG-4	5.88	4.63	1.08	0.17	2		2.5	M	0.89	0.96
WLTG-5	11.64	8.84	2.4	0.4	1.17	0.78	1.83	L	0.8	0.91
WLTG-6	119.19	87.94	25.53	5.71	0.58	0.81	0.88	Н	0.87	0.99
$Pre-wltga$	28.53	20.5	6.88	1.15	1.23	0.8	1.85		0.77	0.97

LREY light rare earth elements including La, Ce, Pr, Nd, and Sm, MREY medium rare earth elements including Eu, Gd, Tb, Dy, and Y, HREY heavy rare earth elements including Ho, Er, Tm, Yb, and Lu, REY are normalized by Upper Continental Crust reported by Taylor and McLennan ([1985\)](#page-4-0), L light REY enrichment type, M medium REY enrichment type, H heavy REY enrichment type, $\delta Eu = Eu_N/(Sm_N \times Gd_N)^{1/2}$; δ Ce = Ce_N/(La_N × Pr_N)^{1/2}

^a Data from Dai et al. [\(2012a](#page-4-0))

and world coals are normalized to the UCC (Fig. 3). The distribution curves of REY averages for both Chinese and world coals are similar to those of the UCC, with exception of Tm.

Based on the three-fold classification (Seredin and Dai [2012\)](#page-4-0), three enrichment types are identified in the coal samples present in this study, including Light-type (light-REY; La_N/Lu_N > 1), Medium-type(medium-REY; La_N/ $Sm_N < 1$, $Gd_N/Lu_N > 1$), and Heavy-type (heavy-REY; $\text{La}_{\text{N}}/\text{Lu}_{\text{N}}$ < 1). The REY compositional characteristics in coal samples WLTG-1 and WLTG-5 (Fig. 4a, c) are characterized by Light-type. The Light-type compositional characteristics of REY in coals are usually caused by the input from terrigenous materials in the sediment source region during peat accumulation (Dai et al. [2012a](#page-4-0)). The LREY-type in the coals is probably attributed to the occurrence of light REY-rich magmatic rocks (granite or carbonatite) and deep weathering products from granite (Seredin and Dai [2012](#page-4-0)). The Medium-type compositional characteristics of REY in WLTG-4 (Fig. 4c), with a relatively strong pronounced Eu-minimum, are due to acid natural waters (Johanneson and Zhou [1997](#page-4-0)) and are probably derived from acid hydrothermal solutions (Seredin and Dai [2012\)](#page-4-0). The Heavy-type compositional characteristics are shown in Fig. 4b, d. Because of the wide spread of HREY-enriched natural water that may circulate in coalbearing strata, the precipitated REY-enriched matter lead to the higher HREY content in the coal samples (e.g., Seredin [2001](#page-4-0); Seredin and Dai [2012](#page-4-0)). These waters may be high-temperature (>500 °C) volcanogenic solutions (Rybin et al. [2003](#page-4-0)), alkaline terrestrial fluids (Johanneson and Zhou [1997\)](#page-4-0), low-temperature (130 °C) alkaline hy-drothermal waters (Michard and Albarède [1986\)](#page-4-0), some high $pCO₂$ cold mineral solutions (Shand et al. 2005), and seawater (Elderfield and Greaves [1982](#page-4-0)).

Fig. 3 REY plots for average values in Wulantuga (WLTG), Chinese coals (CC) and world coals (WC). Data of Chinese coals and world coals are from Dai et al. [\(2012b](#page-4-0)) and Ketris and Yudovich [\(2009](#page-4-0)) respectively. Normalized by average REY content of Upper Continental Crust (UCC) (Taylor and McLennan [1985\)](#page-4-0)

Fig. 4 REY compositional characteristics in the No. 6 coal in the Wulantuga, Inner Mongolia, China

5 Conclusions

- (1) The concentration of REY in study area varies from 5.88 to 119.19 μ g/g, with an average content of $61.85 \mu g/g$, similar to the average values for world low-rank coals.
- (2) The light-REY and heavy-REY are the main REYenrichment type of the Ge-poor coal from the Wulantuga ore deposit. The top and bottom parts of the coal seam have a higher content of REY than that of the middle part.

(3) The occurrence of REY in the Ge-poor coal of Wulantuga may be associated mainly with terrigenous materials input and natural waters enriched in yttrium and rare earth.

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