



The slope creep law for a soft rock in an open-pit mine in the Gobi region of Xinjiang, China

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Abstract The lithology of the strata in the Gobi region of the Xinjiang autonomous region of China is mainly composed of mudstone, silty mudstone, and other soft rocks. Because of the low strength of the rock mass and the serious effects of physical weathering in this area, the slope stability in open-pit mines is poor, and creep deformation and instability can readily occur. Taking the Dananhu No. 2 open-pit mine as a typical example, the creep test of a mudstone sample under different stress levels was studied. Then, based on a bottom friction experiment and a FLAC3D numerical simulation, the deformation and failure processes of the slope were analyzed. The stress–displacement curve and the displacement–time curve for the monitoring points were plotted to obtain the relationship between the stress and displacement for the slope of the soft rock. The results showed that the long-term strength of the mudstone was between 8.0 and 8.8 MPa, and that stable creep occurred when the slope was under low stress. The potential failure mode for this type of slope is that the front edge creeps along the weak layer and then a crack is formed at the trailing edge of the slope. When the crack penetrates the weak layer, cutting bedding and bedding sliding occur. The deformation process of the stable creep slope includes an initial deformation stage, an initial creep stage, a constant velocity creep stage and a deceleration creep stage.

Keywords Gobi region · Soft rock slope · Numerical simulation · Creep deformation · Creep experiment

1 Introduction

Soft rock is widely distributed in China. Given a relative depletion of shallow resources in the country, the depth and scale of open-pit mining have been increasing recently

such that soft rock slopes are exposed and soft rock landslides may occur. The deformation and instability of soft rock slopes are usually accompanied by the occurrence of creep. Due to the complex and diverse nature of the engineering geological conditions, together with the formation, deformation, instability and other induced factors, the creep slope has become an active research field in the mining industry in recent years (Yang et al. 2003; Fan and Gao 2007; Tan et al. 2010; Xu 2017).

For a sustained period, researchers in China have studied creep deformation and the instability mechanism (Wang 2009; Wang et al. 2010) in soft rock slopes, and also the failure mechanisms (Li et al. 1998; Wang et al. 2013) and reinforcement and treatment measures (Wang et al. 2007). These studies have provided a theoretical basis for the prevention and control of failures of soft rock slopes. However, there are few studies on the creep of soft rock slopes which consist of multi-rock and multi-coal seams in

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open-pit mines in the western regions of China such as in the Xinjiang and the Inner Mongolia autonomous regions.

The study site selected for research was based on the east slope of the Dananhu No. 2 open-pit coal mine in the Xinjiang autonomous region. In combination with the topography and engineering profile of the slope of the mining area, the creep deformation law for the soft rock slope was examined using a simulation test and numerical simulation of the bottom friction, the aim being to provide a theoretical basis for the safe and efficient operation of open-pit mines.

2 Geological conditions of mining area

The Dananhu No. 2 coal mine is located in the northwest of the Dananhu mining area in Xinjiang province. The exposed strata in the coal mine was formed in the Mesozoic Jurassic, the Neogene and the Quaternary. The surface of the mining area is composed of fine sand, silt, loess, gravel and clay of the Quaternary and Neogene, which collectively consist of loose rock sections with poor stability and brittleness. Most of the strata below the loose layer consist mainly of mudstone, sandy mudstone and coal, followed by sandstone. The strength of rock of the east slope is less than 20 MPa. It is known from the definition of soft rock (Cai et al. 2002; He et al. 2012) that the rock belongs to the soft rock category. The Jurassic soft rock is composed mainly of weak surfaces. The argillaceous rock structure is weak and cracked, and the hydrological properties are poor. The rock structure is readily softened after water–rock interaction such that the shear strength is greatly reduced. In the dry and cold climate of the Xinjiang region, due to the rapid changes in temperature, the rock mass is subjected to strong physical weathering after exposure from mining, resulting in secondary fractures, and significant decreases in rock stability.

3 Rock mechanical strength and creep experiments

A large amount of monitoring data concerning slope engineering studies in open-pit mines show that the mechanical response of the formation during mining often displays a clear time effect, and the time effect of the deformation will affect the deformation of the slope, which is caused by the rheological properties of the rock layer. The rheology of a rock is a characteristic of the viscous fluid flow that is manifested by deformation over time. (He 2004; Ding et al. 2005; Wang et al. 2009). The Dananhu No. 2 mine is dominated by soft rock; thus the rheological properties are relatively straightforward and predictable.

Therefore, it is of great interest to study the variation of rock and soil strength with time in the Dananhu No. 2 open-pit mine.

3.1 Experimental scheme

Combined with the creep properties of soft rock, the uniaxial compression rheological test was performed. The apparatus for performing the creep experiment was the MTS electro-hydraulic servo system shown in Fig. 1. Before the creep test, the rock uniaxial compressive strength test of the rock sample was undertaken to obtain the instantaneous compressive strength of the rock, and this test was used as the basis for estimating the applied load.

The creep curve of the soft rock was obtained by hierarchical loading with the experiment being divided into 7 stages. The stress interval for levels 1–4 was 1.6 MPa. From the 5th level, the interval was 0.8 MPa, and the loading stresses were 1.6, 3.2, 4.8, 6.4, 7.2, 8.0 and 8.8 MPa, respectively. The initial displacement value was obtained at the time of the experiment loading, and for the later stages of the experiment, the displacement values were recorded at 20-min intervals.

3.2 Analysis of results

The uniaxial creep curve for the mudstone in Fig. 2 shows: the rock-only attenuated creep and the constant velocity creep under the loading stress conditions 1–6 indicate that



Fig. 1 Experimental system

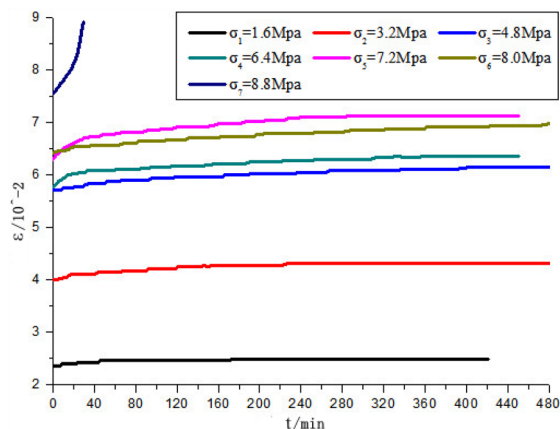


Fig. 2 Uniaxial creep tests for mudstone

the initial stage of creep is small when the constant loading on the rock is small. At the faster speeds, the rock was instantaneously deformed. With increase in time, the deformation of the sample tended to stabilize. When the loading stress reached 8.8 MPa, the strain rate of the rock changed rapidly and three stages of creep developed in a short time. Internal cracks developed rapidly, the deformation intensified and the final damage indicated that the creep became unstable.

Therefore, it can be judged that the long-term strength of the soft rock was between 8.0 and 8.8 MPa. Under the stress levels of 1–6, the stress was less than the long-term strength, the rock first undergoing instantaneous creep, and then the strain of the rock increased continuously, the rate of deformation gradually decreasing, eventually becoming constant, indicating that the rock creep had terminated, and the creep had stabilized. When the loading stress was greater than the long-term strength of the soft rock, the strain of the rock increased sharply in a short period of time, and cracks were generated in the sample, eventually resulting in breakage and corresponding to unstable creep.

4 Analysis of creep deformation in a soft rock excavation slope

4.1 Numerical analysis

4.1.1 Establishment of slope model

The FLAC3D numerical simulation software was used to establish the geological model for the east slope of the open-pit mine. The front and back of the model were constrained by the Y -direction displacement. The left and right directions of the model were constrained by the X -direction displacement. The bottom of the model was constrained by the Z -direction displacement which formed the displacement

boundary and maintained the balance of the system. This simulation used the Mohr–Coulomb constitutive model (Peng 2008; Liu et al 2010) for calculation and analysis; the model of the slope is shown in Fig. 3.

4.1.2 Mechanical parameters of the rock

According to the on-site geological exploration data, the lithology of the top and bottom of each coal seam was mainly mudstone, silty mudstone and sandstone, which readily formed a weak layer. The weak layer in the 16-coal seam was mainly mudstone and the weak layers in the 18-, 21-, 22-, 25- and 28-coal seams were mostly silty mudstone which were softened when exposed to water. Based on previous experimental data, the physical and mechanical parameters of the rock (soil) of the stratum are given in Table 1.

4.1.3 Simulation results

Figures 4, 5, 6, 7 present the numerical simulation results of the east slope. The deformation history of the current slope under gravity stress was studied by analyzing the total displacement cloud map, the displacement vector cloud map, the shear stress cloud map and the plastic zone cloud map.

It can be seen from the total displacement cloud diagram (Fig. 4) and the displacement vector cloud diagram (Fig. 5) that the top surface of the slope has the largest deformation, the displacement vector at the top of the slope is vertically downward, the middle is inclined downward, and the mid-lower slope has a horizontal shearing trend. These data indicate that the displacement deformation mainly consisted of subsidence. Due to the excavation and unloading, the slope had a tendency to slide along the weak layer. Figure 6 shows that the internal stress distribution of the slope was uniform, and there was no obvious stress concentration; the slope can, therefore, be considered to have been in a stable state; From the plastic map of the slope as shown in Fig. 7, the coal seam and the soft rock of the slope in the weak layer were affected mainly by the shear stress, and also the plastic deformation zone appeared. The slope top was deformed plastically due to the tensile stress and it did not form a weak joint with the internal mudstone. The main deformation of the slope was controlled by the weak layer of mudstone, and the top and bottom parts of

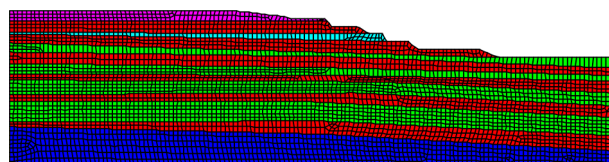
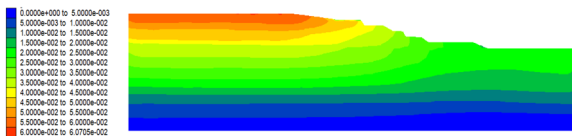
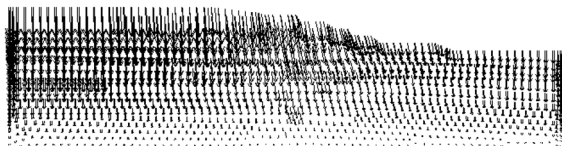
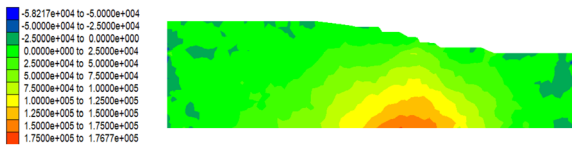
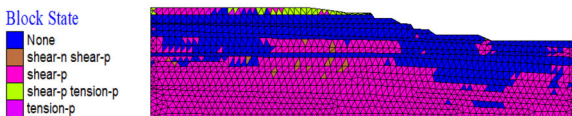


Fig. 3 Present slope model

Table 1 Mechanical parameters of the rock

Rock mass	Bulk, K (MPa)	Shear, G (MPa)	Cohesion, C (kPa)	Friction angle, Φ ($^{\circ}$)
Gravel sand	25	11.5	4	26.8
Mudstone	400	230	139	18.6
16-Coal	330	160	109	17.7
Silty mudstone	280	120	85	17.9
Sandstone	570	340	245	19.7

**Fig. 4** Total displacement**Fig. 5** Displacement vector**Fig. 6** Shear stress**Fig. 7** Plastic zone

the coal seam were the main control sliding surfaces of the slope.

4.2 Monitoring point displacement analysis

To study the variation of displacement of the slope under excavation and unloading and gravity stress, monitoring points were arranged at the top, the middle surface and the bottom of the slope. The main monitoring points selected for analysis were 3–2, 3–4, 3–6, 3–7 and 3–8 with the positions of the monitoring points being indicated in Fig. 8.

From inspection of Fig. 9, it can be seen that the displacement of the monitoring points changed on a stepwise

basis. The displacement can be divided into three stages, namely, the slope was affected by the excavation and the unloading and the gravity stress; also the displacement of the initial monitoring point changed slightly, going below the 0 level. Under the influence of excavation, unloading and gravity stress, the displacement increased sharply. After stress redistribution and structural adjustment, the displacement deformation rate slowed down and gradually stabilized, eventually reaching a stable value. Within this, the displacement deformation of monitoring point 3–6 was the maximum, which was due to the presence of the weak coal seam layer and rock below the existing surface of the slope, and which was affected by the shearing displacement of the weak layer of soft rock. This was followed by 3–7 and 3–4; the displacement at 3–8 and 3–2 was small. The lower part of the bench in the bottom of the mining pit was subjected mainly to squeeze displacement caused by the shearing of the soft rock weak layer of the slope. Compared with the area below the top of the slope, deformation was greater.

5 Similar simulation test for slope of soft rock

5.1 Experimental principle and design

The bottom friction experiment was based on the similarity principle (Gu 1995; Wang et al. 2018). When various conditions are similar (geometry, stress, deformation, etc.), friction may be used as a proxy for gravity and then the deformation and failure rules for rock and soil under the action of gravity can be studied. Based on this similarity, the slope similarity model was constructed.

The bottom friction experimental device is illustrated in Fig. 10. A similar material model 2 was placed on the drive belt (steel sheet) 1 which was propelled by an electric motor (manual operation also possible). Thus, when the belt moves in the direction shown in the drawing, the model and the belt will generate friction due to the blocking of the cross bar 4. In this way, the bottom friction model experiment uses friction to simulate the gravity of the rock formation. This happens because the friction

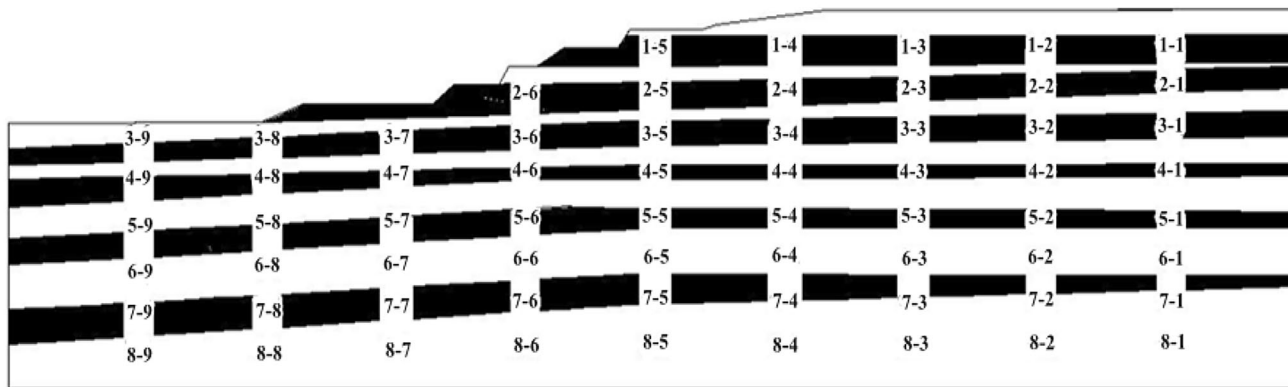


Fig. 8 Layout of the slope displacement monitoring points

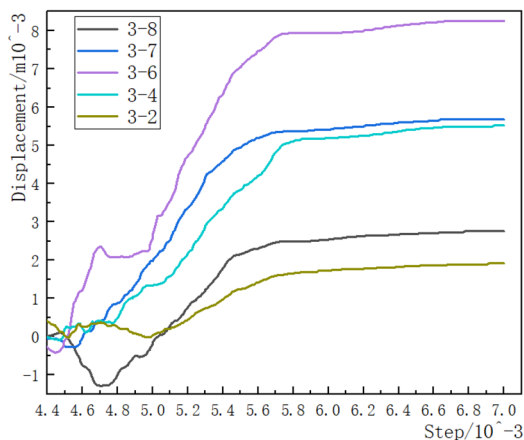


Fig. 9 Displacement variation curves for the monitoring points

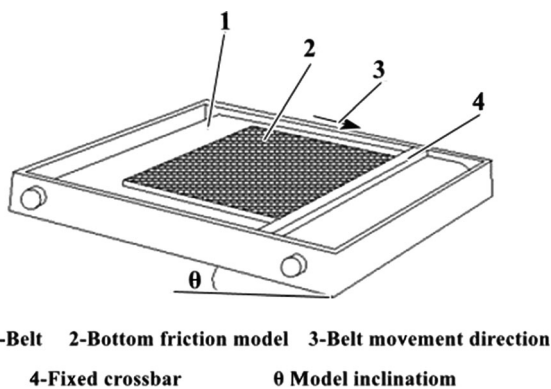


Fig. 10 Bottom friction test device

applied to the surface of the model is in direct contact with the belt. The point where gravity acts should be applied to the intermediate layer of the model along the thickness direction. The thickness of the bottom friction model is generally kept small to reduce the error caused by the frictional force and the point of gravity, thus ensuring the validity of the test.

Regarding the simulation test, the frame size of the bench was $1\text{ m} \times 0.9\text{ m}$, fine sand was used as the raw material, and lime and gypsum were used as cementing materials (various proportions) to simulate the slope model. To facilitate the observation of the deformation and the statistical treatment of the results, the model was vectorized and a grid of dimensions $100\text{ mm} \times 30\text{ mm}$ grid was marked out on the surface of the model.

5.2 Experimental process analysis

The evolution diagram for the deformation of the slope under gravity stress and excavation and unloading are presented in Fig. 11. For a run of 120 s, two cracks corresponding to S1 and S2 appeared, which developed at the bottom line of the slope and stretched downwards. When running to 170 s, vertical tension cracks corresponding to S3 appeared along the bottom line of the step. Due to the cutting action of the crack and the weak layer of the coal seam–rock interface, when running to 190 s, three tensile cracks appeared at the top of the slope, that is, S4, and the downward expansion caused partial destruction of the 21-coal seam. The other steps showed successive cracks. After running to 210 s, the slope was decomposed into multiple blocks from the top to the bottom of the 21st seam, and along the coal seam. The bottom plate was displaced, and the 22-coal seam was also affected by the corresponding fracturing, and there was a tendency to slide.

5.3 Monitoring point displacement analysis

To study the deformation law and the damage caused by excavation and unloading and gravity stress of the slope due to creep, the horizontal displacement of the same row of monitoring points was analyzed statistically. The monitoring points selected were 3-2, 3-4, 3-6, 3-7 and 3-8 and the results are presented in Fig. 12.

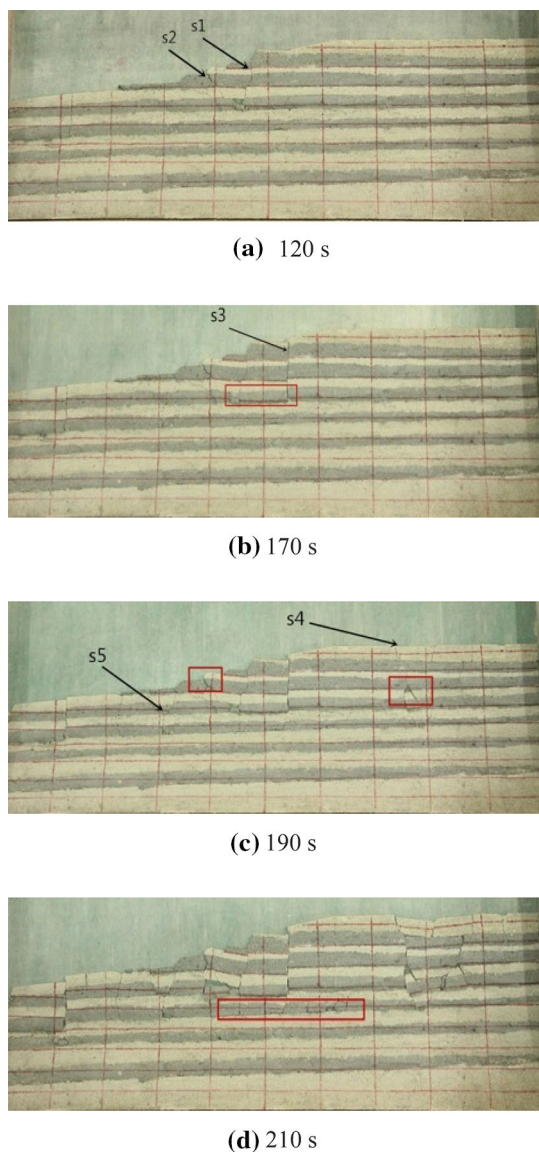


Fig. 11 The slope deformation process

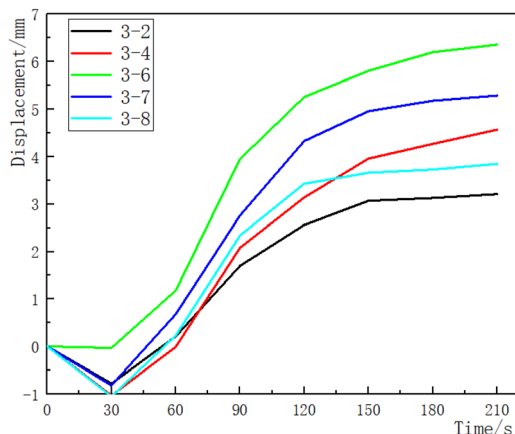


Fig. 12 Displacement variation curves for monitoring points

From the time–displacement curve, the displacement trend was an “S” type curve, the initial displacement fluctuating above 0, and then increasing rapidly. After a period of time, the displacement began to slow down and gradually reached a specific value, which was consistent with the trend from the numerical simulation. The displacement variations at monitoring point 3–6 was maximum, followed by 3–7 and 3–4; the displacement variations for monitoring point 3–2 at the trailing edge of the slope and for 3–8 at the leading edge of the slope were small. The order of the change in the total displacement was as follows: the middle of the slope > the front edge of the slope > the trailing edge of the slope.

These findings indicated that after the excavation, the slope was affected by gravity stress and unloading, such that creep deformation occurred inside the slope, and micro-cracks S1 and S2 were generated inside the slope, and the cracks gradually expanded to the surroundings over time, and the displacement sharply increased. The monitoring points 3–6, 3–4, 3–7 increased gradually as a result of the expansion of the cracks, and 3–6 was most affected. The internal creep of the slope did not extend to the surface of slope, so the displacement at point 3–8 was smaller than that in the slope, but it was still affected by the extrusion displacement of the weak layer of soft rock in the slope. The subsidence deformation of 3–2 was mainly dominated by gravity stress, which had the least impact.

Based on the numerical simulation and the bottom friction similar simulation test results, it has been shown that the cracks of the slope were connected to the lower weak layer by the stress concentration of the east bench, and where the weak layer sheared and underwent creep. The slope relied mainly on its own gravity to generate the sliding force. Due to the horizontal tensile stress of the slope, tensile deformation occurred at the top of the slope, and a tensile failure zone formed. The shear extrusion deformation occurred at the interface between the weak rock and the coal seam. The horizontal displacement had a large amount of deformation, and the rock mass on the upper surface of the slope produced vertical tensile micro-cracks. Lateral deformation and expansion occurred in the weak layer of mudstone inside the slope. With development of the tensile fracture zone, the rock mass on the upper surface of the slope generated vertical tensile micro-cracks. The tensile cracks gradually penetrated into the horizontal slip surface, which led to further development of the slope deformation, and there was the tendency for sliding to occur along the bottom of the slope. Therefore, the deformation of the eastern slope was mainly controlled by the weak layer of mudstone, and the 22-coalseam roof was the main sliding surface of the slope. The deformation was dominated by the extension of the vertical tensile crack

and the shearing of the soft layer. The slope slide was of the creep-cracking mode.

6 Analysis of the slope creep law for soft rock

6.1 Relationship between stress variation and displacement for the excavation slope of soft rock

According to the uniaxial compression creep experiment for mudstone and the slope deformation in the similar simulation experiment, the shear creep deformation of the slope for layered soft rock has great influence on the slope stability, thus it is necessary to study the area of the creep (Fan and Gao 2007; Fan et al. 2010). Through analysis of the FLAC3D numerical simulation and the bottom friction test, the stress–displacement curve and the displacement–time curve for the monitoring points were analyzed. Thus, it was possible to establish the corresponding relationship between the stress and strain and the displacement deformation of the slope of the soft rock as shown in Fig. 13. Slope deformation can be divided into four stages: the initial deformation stage, the initial creep stage, the constant velocity deformation stage and the deceleration deformation stage. The analysis further revealed the deformation law for the slope of the soft rock in the Gobi region.

The initial deformation stage (O–A): the slope stress increased rapidly and was in a stage of elastic deformation. Due to the excavation and unloading, the stress in the slope direction increased, and the stress concentration occurred locally at the foot of the eastern slope of the bench, and micro-cracks appeared. After the stress was reduced slightly, the slope entered the creeping stage.

The initial creep stage (A–B): the deformation at this stage was mainly caused by the closure of micro-cracks closure and the structural adjustment. Due to the influence of gravity stress, the lower weak layer was compacted, and the rate of the deformation decreased with time, so that the curve bent in a downward direction. The displacement still increased and was accompanied by small amounts of micro cracks.

The constant velocity deformation stage (B–C): the slope stress continuously adjusted and changed, and the internal damage effect of the slope was enhanced, so that the micro cracks began to expand, increasing continuously and evenly in a downward direction; this made the horizontal displacement of the creep zone increase. Therefore, compared with the A–B segment, the deformation rate increased.

The deceleration deformation stage I (C–D): Point C was the turning point. After point C, the internal structure of the slope re-adjusted, the plastic deformation of the slope weakened and the velocity of the creeping movement slowed down. The macroscopic performance showed that the displacement rate of the slope decreased compared with the constant velocity deformation stage.

The deceleration creep stage II (D–E): After point D, the displacement rate was further reduced and tended to zero after point E, and where there was no risk of landslide.

6.2 Slope creep evolution model for soft rock

The study of slope creep for soft rock (Tian et al. 2018) is often based on the instability of the slope; that is, the slope has eventually been destroyed prior to the landslide occurring. The deformation process generally includes a transient deformation phase, an initial deformation phase, a constant velocity deformation phase and an accelerated deformation phase. Due to the excavation and unloading

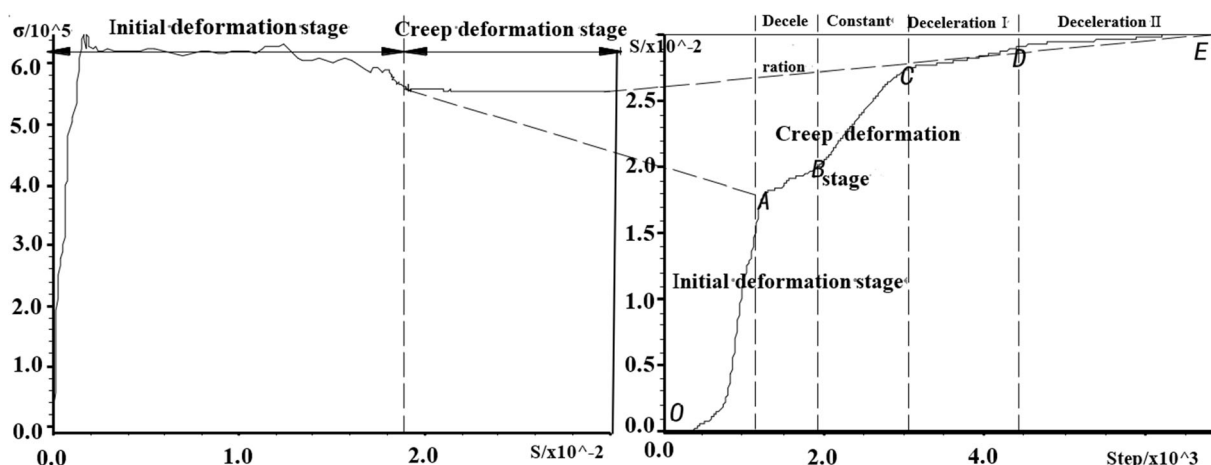


Fig. 13 Deformation process of creep zone

and gravity stress on the slope, the slope has different levels of deformation. When the slope anti-sliding force is greater than the sliding force, the slope can still be in a stable state. Therefore, based on the simulation results and the creep deformation law of the original unstable slope (Li et al. 2013; He et al. 2016), the creep deformation law for a stable slope is proposed, and the original creep model has been improved, obtaining a stable creep slope. The slip evolution model is shown in Fig. 14.

An initial deformation state ($O-A$): This stage corresponds mainly to the elastic deformation stage; and the slope deformation does not enter the creep stage. The effect of the sliding force is mainly related to the structure and is not affected by time.

An initial creep state ($A-B$): From point A , the slope enters the creep state. Due to the compaction and structural adjustment of the rock mass, an elastic effect occurs after deformation, and the creep rate decreases with time, and this is smallest at point B .

A stable creep state ($B-C$): The adjustment of the slope stress induces the crack to increase in a regular manner, and the strain rate remains unchanged at this stage.

A deceleration creep state I ($C-D$): In the process of deformation, point C is the turning point. Given that a landslide failure condition is absent, a plastic deformation occurs gradually and its rate decreases.

A deceleration creep state II ($D-E$): After point D , the deformation decreases further, the deformation rate gradually decreases and tends to zero. After point E , the slope is in a stable state.

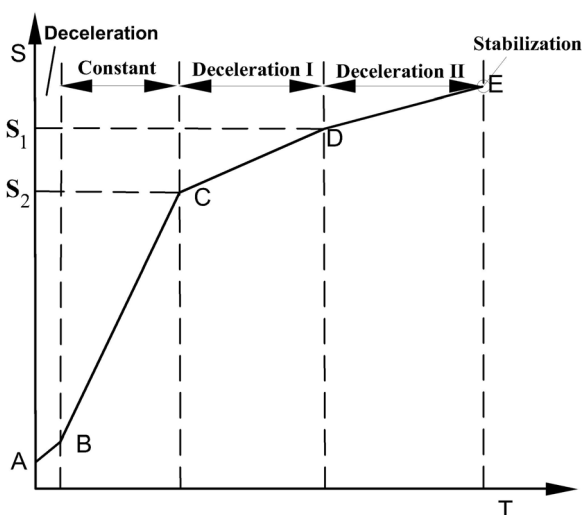


Fig. 14 Schematic of the slope creep evolution model

7 Conclusions

- (1) The long-term strength of mudstone is between 8.0 and 8.8 MPa in the Gobi area of Xinjiang. When the soft rock slope is in a low stress state, the deformation corresponds to stable creep.
- (2) The numerical simulation and bottom friction experiments show that creep deformation of the eastern slope in the Dananhu open-pit mine was mainly controlled by the weak mudstone layer, especially by the material located on the roof and the floor of the coal seam. Also the sliding mode of the slope was creep and tension cracking.
- (3) The stable creep deformation can be divided into four stages: the initial deformation stage, the initial creep stage, the constant velocity deformation stage and the deceleration deformation stage.
- (4) In the process of stable creep deformation, point C in the stress–displacement deformation curve was the critical point. The trend of deformation after point C was the key to determining whether the slope was stable. If the deformation rate after point C decreases, the slope will be in a stable creep condition.

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